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Atmospheric Nitrogen Deposition to Global Forests

Xia, Nan; Du, Enzai; de Vries, Wim

<https://doi.org/10.1016/B978-0-323-91140-5.00018-X>

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# Impacts of nitrogen deposition on soil methane uptake in global forests

Nan Xia<sup>a,b</sup>, Enzai Du<sup>a,b</sup> and Wim de Vries<sup>c,d</sup>

<sup>a</sup>State Key Laboratory of Earth Surface Processes and Resource Ecology, Faculty of Geographical Science, Beijing Normal University, Beijing, China; <sup>b</sup>School of Natural Resources, Faculty of Geographical Science, Beijing Normal University, Beijing, China; <sup>c</sup>Environmental Research, Wageningen University and Research, Wageningen, the Netherlands; <sup>d</sup>Environmental Systems Analysis Group, Wageningen University and Research, Wageningen, the Netherlands

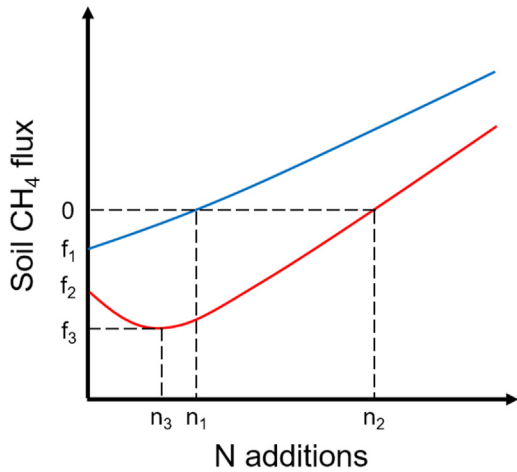
## 1. Introduction

As a well-known greenhouse gas, atmosphere methane (CH<sub>4</sub>) can absorb infrared radiation and has a global warming potential that is 29.8 times higher than CO<sub>2</sub> (Forster et al., 2021; Houghton et al., 2001; Rodhe, 1990). Atmospheric CH<sub>4</sub> concentrations have been increasing rapidly and causing a growing concern of its contribution to climate warming (Dlugokencky et al., 2011; Kirschke et al., 2013). Soil CH<sub>4</sub> can either be produced by methanogens (Le Mer and Roger, 2001; Thauer et al., 2008) or oxidized to CO<sub>2</sub> by methanotrophs (Bosse and Frenzel, 1997; Le Mer and Roger, 2001). Under aerobic conditions, the oxidization of CH<sub>4</sub> generally exceeds its production and thus results in a net soil CH<sub>4</sub> sink (Bodelier and Laanbroek, 2004). Covering nearly one third of the land surface area (Keenan et al., 2015), forest soils are an important sink of atmospheric CH<sub>4</sub> (Dutaur and Verchot, 2007; Le Mer and Roger, 2001). Anthropogenic nitrogen (N) emissions have increased N deposition to global forests (Lamarque et al., 2013; Schwede et al., 2018; Vet et al., 2014), resulting in various ecological consequences (Bobbink et al., 2010; Bowman et al., 2008; de Vries et al., 2011, 2017; Du et al., 2019, 2020a; Gilliam, 2006; Midolo et al., 2019). However, the response of soil CH<sub>4</sub> flux to N deposition is less quantified across global forest biomes in comparison with soil CO<sub>2</sub> fluxes (Bodelier and Steenbergh, 2014; de Vries et al., 2017).

Nitrogen deposition affects both the activities of methanogens and methanotrophs (Bodelier and Laanbroek, 2004; Schnell and King, 1994). In an N-deficient ecosystem, N inputs might release the N limitation of methanotrophic microorganisms and/or the biosynthesis of enzymes involved in methane oxidation and thus benefit soil CH<sub>4</sub> uptake (Bodelier et al., 2000; Bodelier and Laanbroek, 2004; Reay and Nedwell, 2004). In addition, external N

inputs can stimulate tree growth (Du and de Vries, 2018; Schulte-Uebbing and de Vries, 2018; Sonnleitner et al., 2001) and increase evapotranspiration, indirectly favoring microbial CH<sub>4</sub> oxidation in soils. However, excess N inputs may decrease soil CH<sub>4</sub> uptake via a direct inhibition of CH<sub>4</sub> oxidation due to either osmotic stress caused by increased soil inorganic N concentrations or competitive inhibition of the enzyme methane mono-oxygenase by ammonia (Bodelier and Laanbroek, 2004; Schnell and King, 1994; Sitaula et al., 1995) and an indirect effect due to soil acidification and an imbalance of N and P (Veraart et al., 2015). Theoretically, external N additions likely exert nonlinear effects on soil CH<sub>4</sub> uptake in N-limited forests, including an initial stimulation of soil CH<sub>4</sub> uptake and then an inhibition when the N addition rate exceeds a certain threshold (see *red curve* in Fig. 9.1). In non N-limited forests, N additions generally decrease soil CH<sub>4</sub> uptake or increase soil CH<sub>4</sub> emissions (see *blue line* in Fig. 9.1).

Based on an earlier meta-analysis, Liu and Greaver (2009) demonstrated that forest soil CH<sub>4</sub> uptake was significantly reduced by N additions but the analysis was not conducted for each forest biome. In contrast, another meta-analysis indicated an overall neutral effect of N additions on forest soil CH<sub>4</sub> uptake (Aronson and Helliker, 2010). This inconsistency might be caused by mixing experimental results from all forest biomes and neglecting the nonlinear effects of N additions in the meta-analyses (Xia et al., 2020). Previous studies have indicated a poleward increase in N limitation from tropical forest to boreal forest (Du et al., 2020b; Elser et al., 2007), implying distinct responses of soil CH<sub>4</sub> flux to N addition across forest biomes. Therefore, the biome-specific effects of N deposition should be considered separately.



**FIG. 9.1** Conceptual responses of soil CH<sub>4</sub> flux to N additions in N-limited (in red) and non N-limited forests (in blue).  $f_1$  and  $f_2$  indicate the soil CH<sub>4</sub> fluxes under ambient condition (N addition = 0) in non N-limited forests and N-limited forests, respectively.  $n_1$  and  $n_2$  indicate the thresholds of N addition that result in a shift from soil CH<sub>4</sub> sinks to soil CH<sub>4</sub> sources in non N-limited forests and N-limited forests, respectively.  $n_3$  indicates the threshold of N addition that shift from positive to negative effects on soil CH<sub>4</sub> uptake in the N-limited forests.

According to an assessment conducted by the World Meteorological Organization (WMO) Global Atmosphere Watch program (GAW), ambient N deposition rarely exceeded  $60 \text{ kg N ha}^{-1} \text{ yr}^{-1}$  across the globe (Vet et al., 2014). In view of an urban hotspot phenomenon, N deposition is even lower in the widespread natural forests in comparison with urban and rural forests (Du et al., 2022). Current understanding of the effects of N deposition on forest soil CH<sub>4</sub> uptake is mainly derived from manipulated N addition experiments. However, much higher N dosages (e.g.,  $>100 \text{ kg N ha}^{-1} \text{ yr}^{-1}$ ) have been frequently used in existing N addition experiments (Aronson and Helliker, 2010; Bodelier and Steenbergh, 2014; de Vries et al., 2017; Liu and Greaver, 2009; Xia et al., 2020). These experiments might lead to biases when evaluating the effects of actual N deposition on soil CH<sub>4</sub> uptake.

Forest soil CH<sub>4</sub> flux is regulated by several climate factors (Borken et al., 2006; Ni and Groffman, 2018) and soil properties (Benstead and King, 2001; Boeckx et al., 1997; Castro et al., 1995). Compared to methanogens, methanotrophs have a lower optimum temperature and are thus less sensitive to temperature (Borken et al., 2006; Dunfield et al., 1993). Under low temperature conditions (e.g., boreal forest), methanotrophs are likely more active than methanogens, resulting in larger soil CH<sub>4</sub> sinks. Precipitation has a strong control of soil moisture, which not only affects the activity of methanotrophs but also affects the diffusion rates of CH<sub>4</sub> and O<sub>2</sub> molecules (Smith et al., 2003). In soils with low moisture, CH<sub>4</sub> is more easily to diffuse into methanotrophs, which in turn promote a high

oxidation rate. Soil pH affects the physiological functions of methanogens and methanotrophs directly (Benstead and King, 2001) and may interact with other factors (e.g., water stress, terpenes, and ammonium inputs) to limit methanotrophs activity (Amaral et al., 1998a; King, 1997). However, the reported effects of soil pH on soil CH<sub>4</sub> flux vary largely among sites and studies (Hütsch et al., 1994; Saari et al., 2004; Weslien et al., 2009).

Based on a synthesis of experimental results in literature, we evaluated the impacts of N additions on soil CH<sub>4</sub> uptake in global forests. Specifically, the effects of low-level N addition ( $\leq 60 \text{ kg N ha}^{-1} \text{ yr}^{-1}$ ) and high-level N addition ( $>60 \text{ kg N ha}^{-1} \text{ yr}^{-1}$ ) on soil CH<sub>4</sub> flux were assessed separately. Moreover, we explored the key factors affecting the spatial variation in the responses of soil CH<sub>4</sub> flux to N additions, including experimental duration, N addition rate, N addition form, mean annual temperature (MAT), mean annual precipitation (MAP), ambient N deposition, soil pH and forest type. This chapter extends a previous meta-analysis by Xia et al. (2020) on the effects of low and high level N deposition on soil CH<sub>4</sub> uptake in forest biomes, by using an updated literature database and quantifying the response of soil CH<sub>4</sub> uptake per kg N addition for major forest biomes and the global-scale impact of N deposition on forest soil CH<sub>4</sub> uptake.

## 2. Materials and methods

### 2.1 Data set

By conducting a survey via the online library of ISI Web of Science (<http://isiknowledge.com>), Google Scholar (<http://scholar.google.com>) and China National Knowledge Infrastructure (<http://www.cnki.net>), we collected field experimental data on the effect of N addition on soil CH<sub>4</sub> flux in forest ecosystems across the globe. The key words “methane (or CH<sub>4</sub>)”, “nitrogen addition (or nitrogen deposition, nitrogen fertilization)” and “forest” were used. The database recorded the means and standard deviations of soil CH<sub>4</sub> fluxes in control plots and treatment plots, as well as information on site location (longitude and latitude), experimental treatments (the experimental duration, N addition rate, N addition form, i.e., NO<sub>3</sub><sup>-</sup>-N, NH<sub>4</sub><sup>+</sup>-N, NH<sub>4</sub>NO<sub>3</sub>-N or urea), climate (mean annual temperature, MAT, and mean annual precipitation, MAP), ambient N deposition, soil pH and forest type (i.e., boreal, temperate, subtropical or tropical forests) (Table 9.1).

Experimental and climate data were either directly extracted from tables or digitized from figures by using a GetData Graph Digitizer (Version 2.18, <http://www.getdata-graph-digitizer.com>). Soil pH was retrieved from SoilGrids (<http://soilgrids.org>) and ambient N deposition was retrieved from a multiple-model estimate of N deposition (Tan et al., 2018) if it was not reported. In boreal and

**TABLE 9.1** Information on location, climate, forest type, and design of the manipulated N addition experiments in forest ecosystems.

ID	Lon (°E)	Lat (°E)	MAT (°C)	MAP (mm)	Soil pH	N deposition (kg N ha <sup>-1</sup> yr <sup>-1</sup> )	Forest type	N addition form	N addition rate (kg N ha <sup>-1</sup> yr <sup>-1</sup> )	Experimental duration (yr)	References
1	-148.30	64.75	-2.8	284	5.80	0	BF	NH <sub>4</sub> NO <sub>3</sub>	0, 171	4	<a href="#">Gulledge and Schimel (2000)</a>
2	-148.30	64.75	-2.8	284	5.80	0	BF	NH <sub>4</sub> NO <sub>3</sub>	0, 143	4	<a href="#">Gulledge and Schimel (2000)</a>
3	29.37	61.85	3.8	582	4.07	4	BF	NH <sub>4</sub> NO <sub>3</sub>	0, 31	28	<a href="#">Saari et al. (2004)</a>
4	25.62	61.32	3.3	680	3.60	5	BF	NH <sub>4</sub> NO <sub>3</sub>	0, 200	3	<a href="#">Maljanen et al. (2006)</a>
5	125.48	51.37	-2.4	489	5.15	4	BF	NH <sub>4</sub> NO <sub>3</sub>	0, 25, 50, 75	3	<a href="#">Yan et al. (2016b)</a>
6	121.51	50.83	-5.4	580	5.50	9	BF	NH <sub>4</sub> NO <sub>3</sub>	0, 10, 20, 40	0.5	<a href="#">Xu et al. (2014)</a>
7	121.88	50.42	-5.4	500	5.50	10	BF	NH <sub>4</sub> Cl/ KNO <sub>3</sub>	0, 10, 40	1	<a href="#">Gao et al. (2013)</a>
8	129.11	48.12	-0.5	718	5.50	9	TemF	NH <sub>4</sub> NO <sub>3</sub>	0, 50, 100, 150	1	<a href="#">Yan et al. (2016a)</a>
9	128.89	47.18	-0.3	676	5.67	13	TemF	NH <sub>4</sub> NO <sub>3</sub>	0, 20, 40, 80	3	<a href="#">Song et al. (2017)</a>
10	141.33	43.10	9.5	995	4.21	10	TemF	NH <sub>4</sub> NO <sub>3</sub>	0, 50	1	<a href="#">Kim et al. (2012)</a>
11	128.10	42.40	4.1	855	5.90	23	TemF	NH <sub>4</sub> Cl/ KNO <sub>3</sub>	0, 45	0.25	<a href="#">Xu et al. (2011)</a>
12	-85.40	42.40	9.7	890	5.47	6	TemF	NH <sub>4</sub> NO <sub>3</sub>	0, 10	2	<a href="#">Ambus and Robertson (2006)</a>
13	-85.40	42.40	9.7	890	4.63	6	TemF	NH <sub>4</sub> NO <sub>3</sub>	0, 10	2	<a href="#">Ambus and Robertson (2006)</a>
14	127.63	41.70	3.6	745	5.30	11	TemF	Urea	0, 10, 20, 40, 60, 80, 100, 120, 140	0.5	<a href="#">Geng et al. (2017)</a>
15	-80.03	41.60	9.8	1125	4.00	15	TemF	NH <sub>4</sub> NO <sub>3</sub>	0, 100	1	<a href="#">Chan et al. (2005)</a>
16	116.22	39.99	11.6	630	6.72	33	TemF	NH <sub>4</sub> NO <sub>3</sub>	0, 50, 100, 150	1	<a href="#">Wang (2012)</a>
17	-74.58	39.92	12.3	1143	4.80	15	TemF	NH <sub>4</sub> NO <sub>3</sub>	0, 5, 67	0.33	<a href="#">Aronson et al. (2012)</a>
18	112.15	36.69	8.6	662	7.08	31	TemF	NH <sub>4</sub> NO <sub>3</sub>	0, 50, 100, 200, 400	3.33	<a href="#">Yu et al. (2019)</a>
19	118.70	32.18	15.5	1020	4.50	48	STF	NH <sub>4</sub> NO <sub>3</sub>	0, 50, 100, 150	1.5	<a href="#">Hu et al. (2011)</a>
20	110.45	31.60	10.6	1650	5.70	27	STF	NH <sub>4</sub> NO <sub>3</sub>	0, 50	0.42	<a href="#">Pan (2013)</a>
21	119.70	30.23	15.9	1450	6.50	40	STF	NH <sub>4</sub> NO <sub>3</sub>	0, 40, 120	2	<a href="#">He (2019)</a>

*Continued*

**TABLE 9.1** Information on location, climate, forest type, and design of the manipulated N addition experiments in forest ecosystems.—cont'd

ID	Lon (°E)	Lat (°E)	MAT (°C)	MAP (mm)	Soil pH	N deposition (kg N ha <sup>-1</sup> yr <sup>-1</sup> )	Forest type	N addition form	N addition rate (kg N ha <sup>-1</sup> yr <sup>-1</sup> )	Experimental duration (yr)	References
22	115.06	26.74	17.9	1505	4.26	33	STF	NH <sub>4</sub> Cl/ NaNO <sub>3</sub>	0, 40, 120	1	<a href="#">Li et al. (2015)</a>
23	115.07	26.75	18.0	1489	5.50	23	STF	NH <sub>4</sub> NO <sub>3</sub>	0, 50, 100	0.75	<a href="#">Zhang (2013)</a>
24	117.47	26.18	19.4	1700	5.40	36	STF	NH <sub>4</sub> NO <sub>3</sub>	0, 40, 80	6	<a href="#">Wu (2018)</a>
25	112.54	23.18	21.0	1927	4.07	44	STF	NH <sub>4</sub> NO <sub>3</sub>	0, 100	1.5	<a href="#">Gao et al. (2017)</a>
26	112.17	23.17	21.0	1927	3.80	33	STF	NH <sub>4</sub> NO <sub>3</sub>	0, 50, 100, 150	1	<a href="#">Zhang et al. (2008)</a>
27	112.17	23.17	21.0	1927	3.90	33	STF	NH <sub>4</sub> NO <sub>3</sub>	0, 50, 100	1	<a href="#">Zhang et al. (2008)</a>
28	112.17	23.17	21.0	1927	4.00	33	STF	NH <sub>4</sub> NO <sub>3</sub>	0, 50, 100	1	<a href="#">Zhang et al. (2008)</a>
29	112.17	23.17	21.0	1927	5.50	40	STF	NH <sub>4</sub> NO <sub>3</sub>	0, 150	2.66	<a href="#">Zheng et al. (2016)</a>
30	112.17	23.17	21.0	1927	5.50	40	STF	NH <sub>4</sub> NO <sub>3</sub>	0, 150	2.66	<a href="#">Zheng et al. (2016)</a>
31	112.17	23.17	21.0	1927	5.50	40	STF	NH <sub>4</sub> NO <sub>3</sub>	0, 150	2.66	<a href="#">Zhang et al. (2011)</a>
32	112.83	22.57	22.5	1534	3.87	43	STF	NH <sub>4</sub> NO <sub>3</sub>	0, 50, 100	1	<a href="#">Zhang et al. (2012)</a>
33	107.60	22.38	21.8	1200	3.95	22	STF	Urea	0, 48, 167, 334	2	<a href="#">Zhang et al. (2017)</a>
34	-79.83	9.10	27.2	2715	5.10	9	TroF	Urea	0, 125	12	<a href="#">Veldkamp et al. (2013)</a>
35	-82.25	8.75	20.3	5461	4.70	5	TroF	Urea	0, 125	4	<a href="#">Veldkamp et al. (2013)</a>
36	117.39	5.37	25.2	3098	4.70	3	TroF	Urea	0, 100	4	<a href="#">Mori et al. (2017)</a>
37	-79.20	-4.08	19.4	2230	4.20	5	TroF	Urea	0, 50	4.67	<a href="#">Martinson et al. (2020)</a>
38	-79.20	-4.08	15.7	1950	3.70	5	TroF	Urea	0, 50	4.67	<a href="#">Martinson et al. (2020)</a>
39	-79.20	-4.08	9.4	4500	3.80	5	TroF	Urea	0, 50	4.67	<a href="#">Martinson et al. (2020)</a>

Abbreviations: *BF*, boreal forest; *STF*, subtropical forest; *Tem*, Temperate forest; *TroF*, tropical forest.

temperate forests, soil CH<sub>4</sub> fluxes were generally measured during the growing season, while in subtropical and tropical forests, the measurements were conducted the whole year round. During the nongrowing season, temperature of boreal and temperate forests is relatively low and measurements of soil CH<sub>4</sub> fluxes are lacking. Therefore, our analysis ignored the part of soil CH<sub>4</sub> flux in boreal and temperate forests during the nongrowing season.

Overall, our database included experimental results of 74-N addition trials across 39 forests from 30 sites (Fig. 9.2). Specifically, 52 trials were conducted by using NH<sub>4</sub>NO<sub>3</sub> in 28 forests, 17 trials by using urea in eight forests, and five paired trials by using both nitrate (KNO<sub>3</sub> or NaNO<sub>3</sub>) and ammonium (NH<sub>4</sub>Cl) based N forms in three forests (Table 9.1). For the paired trials, the rates of soil CH<sub>4</sub> flux were averaged for each N dosage and used as one treatment for further analysis. Finally, there were 12, 28, 28 and six fertilization trials from seven boreal forests, 11 temperate forests, 15 subtropical forests and six tropical forests, respectively (Table 9.1). However, the tropical forest studied by Martinson et al. (2020) was a mountain rain forest with low soil mineral N availability and a closed soil N cycle, which was different from most tropical forests. Therefore, we considered this study as a special case and excluded the 3 N addition trials. Afterward, data of tropical forests were combined with those of subtropical forests for further analysis since there were too few trials in tropical forests for a robust statistical analysis (n = 4).

## 2.2 Statistical analysis

To qualify the effect of N additions on soil CH<sub>4</sub> flux, we defined a response ratio (*RR*, g CH<sub>4</sub> kg<sup>-1</sup> N) as given in Eq. (9.1),

$$RR = \frac{(Flux_t - Flux_c)}{N_{add}} \times c \quad (9.1)$$

where *Flux<sub>t</sub>* and *Flux<sub>c</sub>* indicate the mean soil CH<sub>4</sub> fluxes in the treatment plots and control plots (mg CH<sub>4</sub> m<sup>-2</sup> h<sup>-1</sup>), *N<sub>add</sub>* indicates the N addition for each treatment (kg N ha<sup>-1</sup> yr<sup>-1</sup>), and the constant *c* is a unit correction factor from mg CH<sub>4</sub> m<sup>-2</sup> h<sup>-1</sup> to g CH<sub>4</sub> ha<sup>-1</sup> yr<sup>-1</sup>. In subtropical and tropical forests, the unit correction factor (*c* = 87,600) was calculated based on the duration of a whole year. The unit correction factors (*c* = 29,200 for boreal forest and 51,100 for temperate forest) were estimated using durations of 4 months (mid-May to mid-September) for boreal forest and 7 months (April to October) for temperate forest (Piao et al., 2007) while neglecting the remaining time of the year. A positive value of the response ratio indicates a reduction in soil CH<sub>4</sub> uptake or an increase in soil CH<sub>4</sub> emission due to N addition, while a negative value indicates an opposite effect.

The variance (*v<sub>i</sub>*) of the response ratio is estimated according to Eq. (9.2) as given below (Schulte-Uebbing and de Vries, 2018),

$$v_i = \left[ \frac{n_t + n_c}{n_t \times n_c} + \frac{(n_t - 1) \times sd_t^2 + (n_c - 1) \times sd_c^2}{n_t + n_c - 2} \right] \times \left( \frac{c}{N_{add}} \right) \quad (9.2)$$

where *sd<sub>t</sub>*, *sd<sub>c</sub>*, *n<sub>t</sub>*, *n<sub>c</sub>* indicate the standard deviation and number of replicates for the treatment plots and control plots, respectively. Using a random effect model (Borenstein et al., 2009), we estimated the mean response ratios and corresponding standard errors for the forest biomes with weight (*w<sub>i</sub>*) equal to the inverse of the sum of the within-study variance (*v<sub>i</sub>*, due to sampling error) and between-study variance (*τ<sup>2</sup>*, due to variations in experimental conditions) according to Eq. (9.3),

$$w_i = \frac{1}{v_i + \tau^2} \quad (9.3)$$

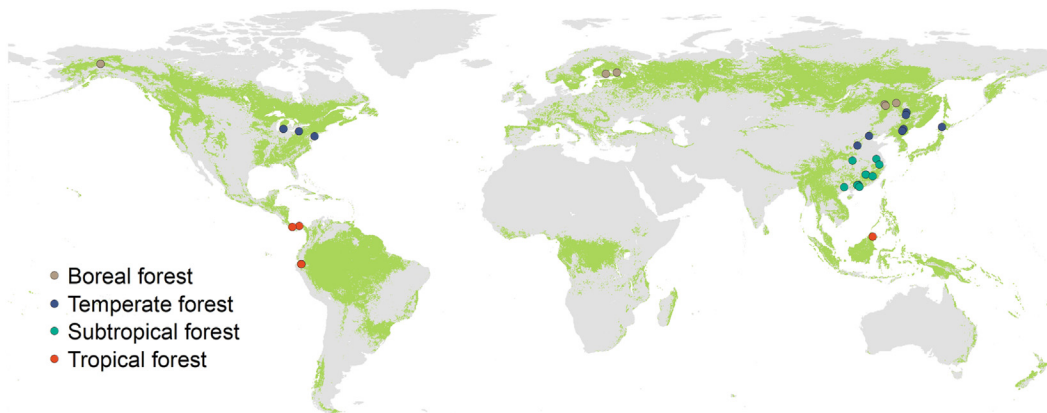


FIG. 9.2 Geographical distribution of nitrogen addition experiments from 39 forests across 30 sites. Green shadows indicate the distribution of global forest.

We conducted one-way ANOVA analysis and a post-hoc test to compare the response ratio of soil CH<sub>4</sub> flux to N addition among boreal, temperate, as well as subtropical and tropical forests. In addition, we used the Student's t test to compare the response ratio between low-level N addition ( $\leq 60 \text{ kg N ha}^{-1} \text{ yr}^{-1}$ ) and high-level N addition ( $> 60 \text{ kg N ha}^{-1} \text{ yr}^{-1}$ ) in different forest biomes respectively. The low-level N addition is in the range of N deposition over the world (Vet et al., 2014). Moreover, biome-scale effects of N deposition on soil CH<sub>4</sub> sink were estimated based on the response ratio of soil CH<sub>4</sub> flux to low-level N addition, mean forest-specific total N deposition (base year 2010) (Schwede et al., 2018) and the area of each forest biomes (base year 2010) (Keenan et al., 2015).

Then we used the “*glmulti*” package to conduct a model selection analysis for the potential driving factors affecting the variation of response ratios of soil CH<sub>4</sub> flux to low-level N addition ( $\leq 60 \text{ kg N ha}^{-1} \text{ yr}^{-1}$ ) based on the corrected Akaike information criterion (Calcagno and de Mazancourt, 2010). The relative importance value of each driving factor was estimated as the sum of the Akaike weights for the models in which the variable appeared. A cut-off relative importance value of 0.8 was used to differentiate between the important and unimportant variables (Calcagno and de Mazancourt, 2010). Conditional regression analysis (Breheny and Burchett, 2013) was conducted to visualize the role of each driving factor on response ratio while holding all the other important variables constant (by default the median for numeric variables). The variance inflation factor (VIF  $< 3$  indicates weak collinearity) was computed to diagnose the multicollinearity of the driving factors for the final model (Zuur et al., 2010). The variance explained by each driving factor was estimated by averaging sequential sums of squares over all orderings of regressors using the “*relaimpo*” package (Groemping, 2006). All analyses were conducted by R software (version 4.0.0,

R Development Core Team, <http://www.r-project.org/>) using a significance level of  $P = .05$ . Values were mean  $\pm$  standard error, if not specially noted.

### 3. Results and discussion

#### 3.1 Variation in ambient soil methane uptake across forest biomes

Forest soils in control plots were significant CH<sub>4</sub> sinks, on average taking up  $0.07 \pm 0.02$ ,  $0.04 \pm 0.01$  and  $0.02 \pm 0.00 \text{ mg CH}_4 \text{ m}^{-2} \text{ h}^{-1}$  in boreal ( $n = 7$ ), temperate ( $n = 11$ ) as well as subtropical and tropical ( $n = 18$ ) forests, respectively. The rate of soil CH<sub>4</sub> uptake in boreal forest was higher than those in temperate ( $P = .08$ ) and subtropical and tropical forests ( $P = .003$ ). The rate of soil CH<sub>4</sub> uptake in temperate forest showed no significant differences from subtropical and tropical forests ( $P = .15$ ). Compared with boreal forest, the lower rates of soil CH<sub>4</sub> uptake in other forest biomes might be attributed to a suppression of CH<sub>4</sub> oxidation by higher temperature (Dutaur and Verchot, 2007), more precipitation (Le Mer and Roger, 2001), higher background N availability (Vet et al., 2014; Schwede et al., 2018; Du and de Vries, 2018), and/or a stronger phosphorus limitation to methanotrophic microorganisms (Veraart et al., 2015).

By multiplying the average rate of soil CH<sub>4</sub> uptake in control plots with the area of each forest biome (Keenan et al., 2015), annual soil CH<sub>4</sub> fluxes were estimated to be  $-2.49 \pm 0.68$ ,  $-1.38 \pm 0.45$  and  $-3.71 \pm 0.56 \text{ Tg CH}_4 \text{ yr}^{-1}$  in global boreal, temperate as well as subtropical and tropical forest biomes, respectively (Table 9.2). Overall, our results showed that global forest soils contributed to a CH<sub>4</sub> sink of  $7.55 \pm 1.68 \text{ Tg CH}_4 \text{ yr}^{-1}$ , accounting for only a half of the estimated annual forest soil CH<sub>4</sub> sink ( $13.9 \text{ Tg CH}_4 \text{ yr}^{-1}$ ) in a model-based study by Murguía-Flores et al. (2018).

**TABLE 9.2** Biome-scale estimates of soil CH<sub>4</sub> fluxes and the effects of nitrogen deposition in global forest biomes.

Forest biome	Area (million ha)	Soil CH <sub>4</sub> flux rate (mg CH <sub>4</sub> m <sup>-2</sup> h <sup>-1</sup> )	Biome CH <sub>4</sub> flux (Tg CH <sub>4</sub> yr <sup>-1</sup> )	Mean N deposition (kg N ha <sup>-1</sup> yr <sup>-1</sup> )	Response ratio (g CH <sub>4</sub> kg <sup>-1</sup> N)	Biome effect (Tg CH <sub>4</sub> yr <sup>-1</sup> )
Boreal	1225	$-0.07 \pm 0.02^b$	$-2.47 \pm 0.68$	1.2	$-11.49 \pm 7.94^b$	$-0.02 \pm 0.01$
Temperate	673	$-0.04 \pm 0.01^{ab}$	$-1.38 \pm 0.45$	7.3	$4.89 \pm 7.57^{ab}$	$0.02 \pm 0.04$
Subtropical and tropical	2118	$-0.02 \pm 0.00^a$	$-3.71 \pm 0.56$	8.3	$10.39 \pm 6.57^a$	$0.18 \pm 0.12$
Total	4016		$-7.55 \pm 1.68$			$0.18 \pm 0.17$

Note: The area of each biome (base year 2010) was derived from FAO Global Forest Resources Assessment (Keenan et al., 2015). Mean N deposition (base year 2010) of each forest biome was derived from Schwede et al. (2018). Different lowercase letters (e.g., a, b) mean significant difference of mean soil CH<sub>4</sub> flux between forest biomes ( $P < .05$ ), and same letters mean no significant difference. The average response ratio of soil CH<sub>4</sub> flux to low-level N addition ( $\leq 60 \text{ kg N ha}^{-1} \text{ yr}^{-1}$ , reasonably indicative to the effects of N deposition) were used to indicate the effect of N deposition.

### 3.2 Biome-specific effects of nitrogen additions on soil methane flux

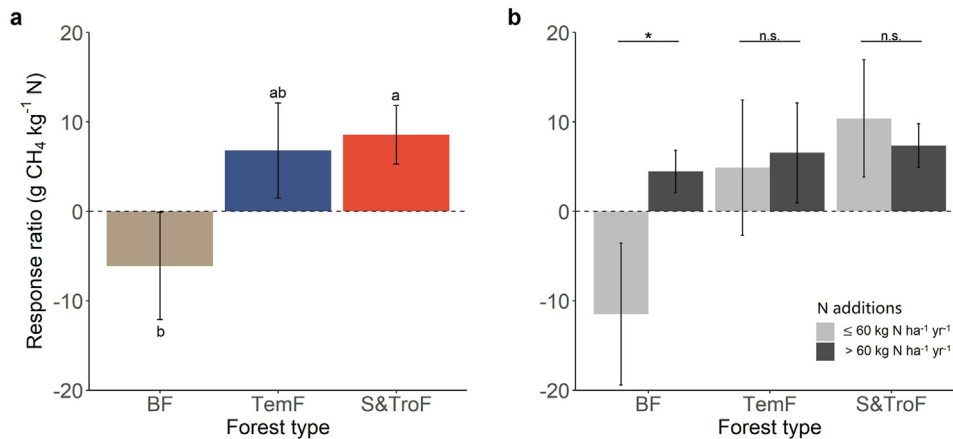
Our results showed that the response of soil CH<sub>4</sub> flux to N addition varied across forest biomes, and the effect of low-level N addition ( $\leq 60 \text{ kg N ha}^{-1} \text{ yr}^{-1}$ ) on soil CH<sub>4</sub> flux was different from that of high-level N addition ( $> 60 \text{ kg N ha}^{-1} \text{ yr}^{-1}$ ) (Fig. 9.3). When both low and high-level were combined, N addition increased soil CH<sub>4</sub> uptake in boreal forest ( $-6.12 \pm 5.99 \text{ g CH}_4 \text{ kg}^{-1} \text{ N}$ ), while it decreased soil CH<sub>4</sub> uptake in temperate forest ( $6.79 \pm 5.31 \text{ g CH}_4 \text{ kg}^{-1} \text{ N}$ ) as well as subtropical and tropical ( $8.55 \pm 3.27 \text{ g CH}_4 \text{ kg}^{-1} \text{ N}$ ) forests (Fig. 9.3a). The response of soil CH<sub>4</sub> flux to low-level N addition, being in the range of atmospheric N deposition, however, was  $-11.49 \pm 7.94 \text{ g CH}_4 \text{ kg}^{-1} \text{ N}$  in boreal forest,  $4.89 \pm 7.57 \text{ g CH}_4 \text{ kg}^{-1} \text{ N}$  in temperate forest and  $10.39 \pm 6.57 \text{ g CH}_4 \text{ kg}^{-1} \text{ N}$  in subtropical and tropical forests, respectively. The difference in response to low-level and high-level N additions was significant in boreal forest ( $P = .04$ ), but insignificant in temperate ( $P = .26$ ) as well as subtropical and tropical ( $P = .31$ ) forests (Fig. 9.3b). Overall, these results indicate a necessity to distinguish among forest biomes and separately assess the effect of low-level N addition from that of high-level N addition in boreal forest. However, many previous meta-analyses did not differentiate the effect of low-level N deposition from that of high-level N addition (Liu and Greaver, 2009; Aronson and Helliker, 2010), which might have overestimated the negative effect of the natural N deposition on forest soil CH<sub>4</sub> sink.

Based on the average response ratio of soil CH<sub>4</sub> flux to low-level N addition ( $\leq 60 \text{ kg N ha}^{-1} \text{ yr}^{-1}$ ), the average N deposition (Schwede et al., 2018) and the area of each forest biome (Keenan et al., 2015), we roughly estimated the biome-scale effects of N deposition on soil CH<sub>4</sub> sink in

global forests. Specifically, N deposition increased CH<sub>4</sub> uptake by  $0.02 \pm 0.01 \text{ Tg CH}_4 \text{ yr}^{-1}$  in boreal forest, while it decreased soil CH<sub>4</sub> uptake by  $0.02 \pm 0.04$  and  $0.18 \pm 0.12 \text{ Tg CH}_4 \text{ yr}^{-1}$  in temperate as well as subtropical and tropical forests, respectively (Table 9.2). Since there is no significant difference of the response ratio between low and high-level N deposition for temperate forest and subtropical/tropical forests, we also estimated the biome-scale effects using mean response ratios for all N addition experiments, being  $6.79$  and  $8.55 \text{ g CH}_4 \text{ kg}^{-1} \text{ N}$  for temperate forest and subtropical/tropical forests, respectively.

Overall, N deposition was estimated to decrease soil CH<sub>4</sub> sink by  $0.18 \pm 0.17 \text{ Tg CH}_4 \text{ yr}^{-1}$  in global forests using the low N deposition responses (Table 9.2). The results showed that when the response ratios to low and high-level N additions were combined in the analysis, CH<sub>4</sub> uptake decreased by  $0.03 \pm 0.03 \text{ Tg CH}_4 \text{ yr}^{-1}$  in temperate forest and  $0.15 \pm 0.06 \text{ Tg CH}_4 \text{ yr}^{-1}$  in subtropical and tropical forests. The overall estimates of biome-scale effects, using the mean response ratios to all N additions, were thus slightly lower than those based on the response ratios to low-level N additions in global temperate, subtropical and tropical forests ( $0.18 \pm 0.09$  and  $0.20 \pm 0.16 \text{ Tg CH}_4 \text{ yr}^{-1}$ , respectively).

Comparison of the mean N deposition induced decrease in soil CH<sub>4</sub> sink of  $0.18 \text{ Tg CH}_4 \text{ yr}^{-1}$  with the total soil CH<sub>4</sub> sink of  $7.55 \text{ Tg CH}_4 \text{ yr}^{-1}$  in global forest (see Table 9.2) implies that this reduction is nearly negligible (2%). Considering that CH<sub>4</sub> is 29.8 times more effective, on a per-unit-mass basis, than CO<sub>2</sub> in absorbing long-wave radiation on a 100-year time horizon (Forster et al., 2021), the warming effect of N deposition induced reduction of soil CH<sub>4</sub> sink in global forest is equivalent to an emission of  $4.68 \text{ Tg CO}_2 \text{ yr}^{-1}$ . This effect on climate warming is negligible in view of the large carbon sink ( $2.4 \text{ Pg C yr}^{-1}$ ) in global established forest (Pan et al., 2011).



**FIG. 9.3** Response ratios of soil CH<sub>4</sub> flux to all N additions (a), low-level ( $\leq 60 \text{ kg N ha}^{-1} \text{ yr}^{-1}$ ) and high-level N additions ( $> 60 \text{ kg N ha}^{-1} \text{ yr}^{-1}$ ) (b) in boreal (BF), temperate (TemF), and subtropical and tropical forests (S&TroF). Error bars are standard errors. Same lowercase letters mean no significant difference and different letters mean significant difference between forest biomes ( $P < .05$ ). The asterisk (\*) indicates a significant difference between low-level and high-level N additions ( $P < .05$ ).



### 3.3 Main drivers of response ratio of soil methane flux to low-level nitrogen addition

The spatial variation in the response ratio of soil CH<sub>4</sub> flux to low-level N additions ( $\leq 60 \text{ kg N ha}^{-1} \text{ yr}^{-1}$ ), reasonably indicative to the effects of actual N deposition, was mainly explained by soil pH and N addition rate (Fig. 9.4a). Conditional regression analysis indicates that the response ratio increased with both higher soil pH (variance explained 17.3%,  $P = .02$ ) (Fig. 9.4b) and N addition rate (variance explained 18.7%,  $P = .01$ ) (Fig. 9.4c). Other factors, including experimental duration, N addition form, MAT, MAP, N deposition and forest type had an unimportant role in shaping the variation of the response ratio of soil CH<sub>4</sub> flux to N addition.

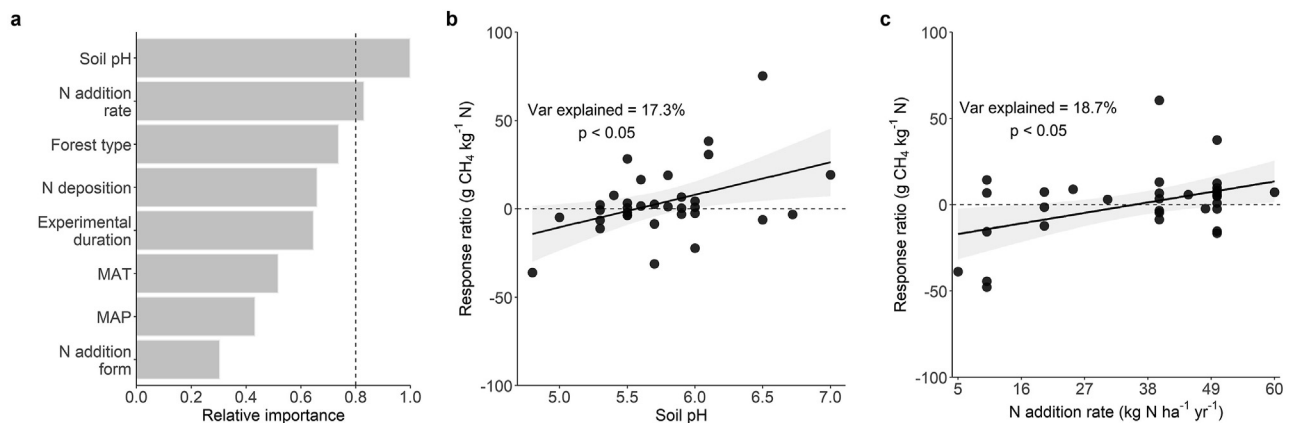
Soil pH plays an important role in the dynamics of soil CH<sub>4</sub> uptake because soil acidity directly affects the physiological functions of organisms, and determines the form and concentration of toxic elements and nutrients (Benstead and King, 2001). Aerobic CH<sub>4</sub> oxidation is mainly driven by methanotrophs in the topsoil (Le Mer and Roger, 2001), which are most active under low soil pH (Amaral et al., 1998b; Saari et al., 2004). Therefore, as soil pH increases, the activity of methanotrophs decreases, thereby inhibiting the absorption/oxidation of soil CH<sub>4</sub>, and changing the response ratio from a negative to a positive value. In addition, soil pH can interact with ammonium inputs to affect the growth and activity of methanotrophs (Amaral et al., 1998a; King, 1997). Nitrification increases with an increase in soil pH, implying a decrease in NH<sub>4</sub> concentration (Dancer et al., 1973; Katyal et al., 1988). However, experimental studies indicate that the ammonium toxicity to soil CH<sub>4</sub> consumption does not vary substantially with pH in moderately acidic soils ( $\text{pH} < 5$ ), but the ammonium toxicity increases with soil pH when  $\text{pH} > 5$  (Benstead and King, 2001). This interaction may also account in part for

the shift from positive to negative effects on soil CH<sub>4</sub> uptake with changes in soil pH (Fig. 9.4b).

Nitrogen addition rate also affected the spatial variation in the response ratio of soil CH<sub>4</sub> flux to N additions. Soil CH<sub>4</sub> uptake decreased with increasing N addition rate, resulting in response ratio shift from negative value to positive value, i.e., a nonlinear effect on soil CH<sub>4</sub> uptake (Fig. 9.4c). Low-level dosages of N addition may release the N limitation of methanotrophic microorganisms and thus increase soil CH<sub>4</sub> uptake in N-deficient conditions (Bodelier et al., 2000; Bodelier and Laanbroek, 2004; Reay and Nedwell, 2004). In contrast, excess N inputs may decrease soil CH<sub>4</sub> uptake via a direct inhibition of CH<sub>4</sub> oxidation due to increased soil inorganic N concentrations (Bodelier and Laanbroek, 2004; Schnell and King, 1994; Sitaula et al., 1995) and an indirect effect due to soil acidification and an imbalance of N and P (Veraart et al., 2015).

## 4. Uncertainties and implications

Uncertainties remain in our analysis of the effects of N deposition on soil CH<sub>4</sub> sink across global forest biomes. First, there were limited numbers of manipulated N addition experiments by using low-level dosage of inorganic N addition, especially in tropical forests. Our search of the literature only gave experimental results for six tropical forests at four sites. Moreover, the three low-level N addition trials that we excluded from our database were all carried out in an atypical tropical forest site (Martinson et al., 2020). Therefore, further efforts are needed to evaluate the effect of N deposition on soil CH<sub>4</sub> flux in tropical forests by conducting N addition experiments with low-level addition of inorganic N. The advantage of the N addition trials in tropical forests is, however, that they were carried out during a whole year. In contrast, the estimates of soil CH<sub>4</sub> uptake and its response to N addition ignored nongrowing season soil CH<sub>4</sub> flux in boreal and temperate



**FIG. 9.4** The relative importance of potential drivers of the spatial variation in the response ratio of soil CH<sub>4</sub> flux to low-level N additions ( $\leq 60 \text{ kg N ha}^{-1} \text{ yr}^{-1}$ ) (a) and the conditional regression plots for soil pH (b) and N addition rate (c). The gray shadows represent 95% confidence intervals for the fitted curves.

forests. This likely underestimated the annual soil CH<sub>4</sub> flux in boreal and temperate forests. If the amount of soil CH<sub>4</sub> sink in nongrowing season is accounted, forest soils may contribute a higher proportion of global soil CH<sub>4</sub> sink, but it will not affect the estimated nearly negligible (2%) N induced impact on this sink.

A recent literature review indicates that the forest soil CH<sub>4</sub> sink in the northern hemisphere has significantly declined during the past three decades and this trend has been attributed to an increase of precipitation (Ni and Groffman, 2018). Our analysis implies that increasing N deposition in developing countries (e.g., China and India) (Abrol et al., 2017; Liu et al., 2013) may partially contribute to the reduction of the soil CH<sub>4</sub> sink. However, N deposition in Europe and the United States has shifted from an increase to a decrease since early or middle 1990s (Du, 2016; Waldner et al., 2014), implying a possible decrease in the effect on soil CH<sub>4</sub> sink since then.

The findings also have implications for improving the performance of process-based models to simulate and quantify the effect of global N deposition on soil CH<sub>4</sub> sinks. So far, the existing models simply account for a negative effect of N inputs on soil CH<sub>4</sub> uptake by involving an inhibition factor (Curry, 2007; Murguía-Flores et al., 2018; Ridgwell et al., 1999). However, our analysis indicates that N addition below a certain threshold level can exert a positive effect on soil CH<sub>4</sub> uptake in boreal forest, implying that existing models most likely overestimate the negative effect of N deposition on forest soil CH<sub>4</sub> sinks, especially in boreal regions. This indicates a necessity of models to separately consider the effect of low-level N addition from that of high-level N addition and distinguish the effects across forest biomes.

## 5. Conclusions and outlook

Our synthesis of experimental results indicates that soil CH<sub>4</sub> flux was significantly affected by N additions in global forest biomes. On a global scale, N deposition was estimated to induce a significant reduction of forest soil CH<sub>4</sub> sink, although it is negligible (2%) compared with the total soil CH<sub>4</sub> sink in global forest. In boreal forest, we found a shift from a positive to a negative effect on soil CH<sub>4</sub> uptake with increasing N addition, while N addition consistently decreased soil CH<sub>4</sub> uptake in temperate as well as subtropical and tropical forests.

Compared with previous assessments based on a meta-analysis approach (Liu and Greaver, 2009; Aronson and Helliker, 2010), our work improves current understanding of actual N deposition effect on soil CH<sub>4</sub> sinks in global forests. It shows that the response of soil CH<sub>4</sub> flux at low-level N additions increases with soil pH and N addition rate, confirming current insights that higher pH values and N availability may reduce soil CH<sub>4</sub> uptake by inhibiting the

activity of methanotrophs and possibly increasing the toxicity of ammonium. Furthermore, the observed biome-specific effects of N deposition on soil CH<sub>4</sub> sinks confirms that the poleward increase in N limitation from tropical forest to boreal forest affects CH<sub>4</sub> oxidation.

Across all forest biomes, the effect of N deposition on soil CH<sub>4</sub> flux is poorly understood in tropical forest due to a lack of experimental studies. This leads to a major uncertainty in the overall effect of N deposition on soil CH<sub>4</sub> sink in global forest. To better understand future trend of soil CH<sub>4</sub> sink in global forest, more experimental and modeling efforts are needed to incorporate the effects of N deposition and other global changing factors.

## Acknowledgments

This work was supported by the Fundamental Research Funds for the Central Universities, the National Natural Science Foundation of China (41877328) and State Key Laboratory of Earth Surface Processes and Resource Ecology (2021-TS-02).

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