

Factors Affecting Mean Flow Velocity of Overland Flow in a Flume

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Introduction

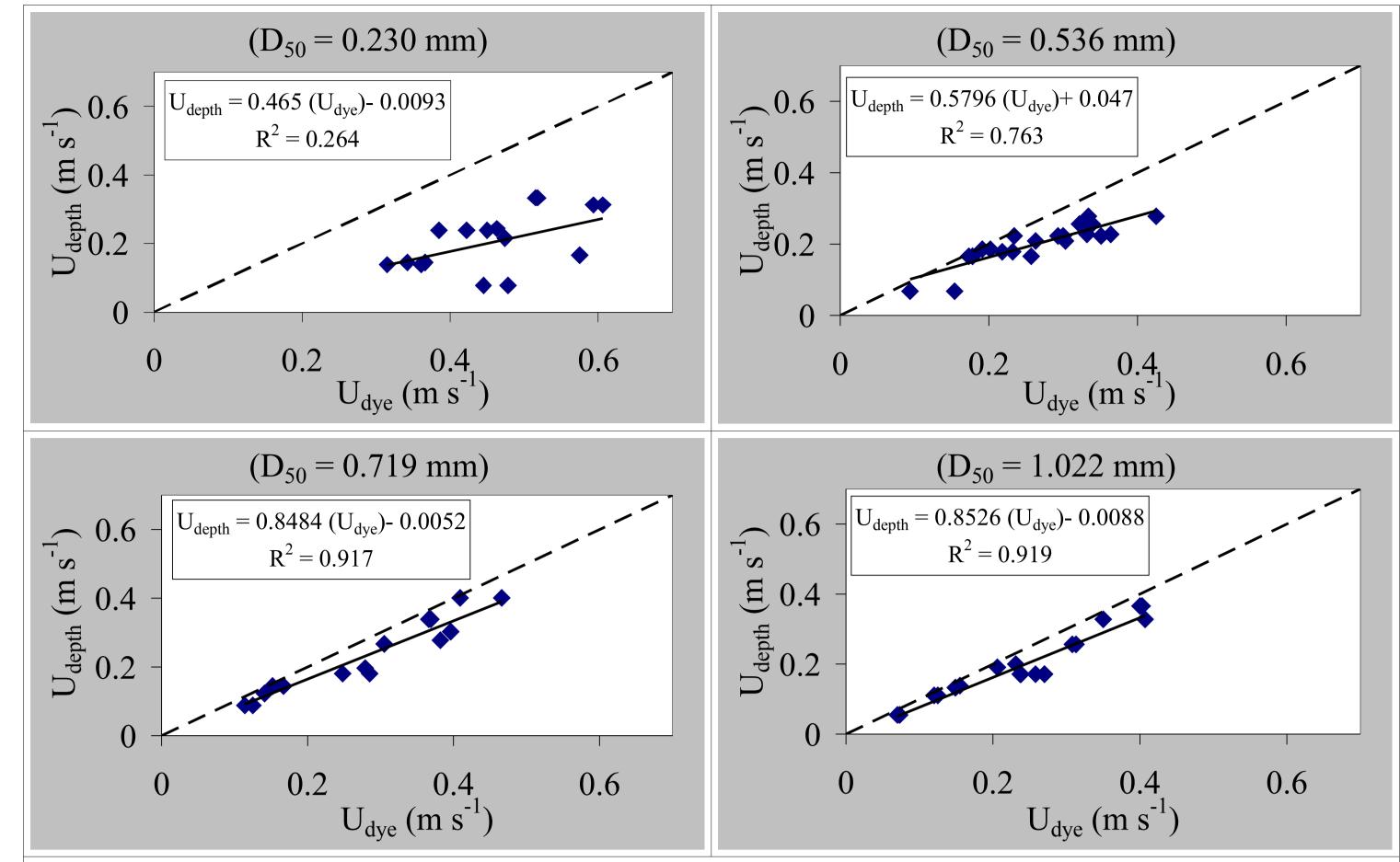
- Mean flow velocity (U_{mean}) of overland flow is determinant hydraulic parameter for hydrological routing and soil erosion modelling.
- Under shallow flow and low discharge estimation is problematic \succ Velocity measurement in general with dye or salt tracer (U_{dve})
 - > Empirical correction factor $\alpha = \frac{U_{mean}}{U_{dya}}$ is applied

 \succ Wider range of α is found in literature for different experimental set-ups

2 Aims

To assess the influence of

- \succ Flow rate (Q)



- > Substrate type (D_{50}), and
- \succ Slope (S)
- on α and on average flow velocity
- To develop a comprehensive equation for estimation of mean flow velocity from dye based estimations

3 Materials and Methods

1 Flume



Figure. 1: Flume with sandy, mobile bed. Point gauges are visible in the foreground.

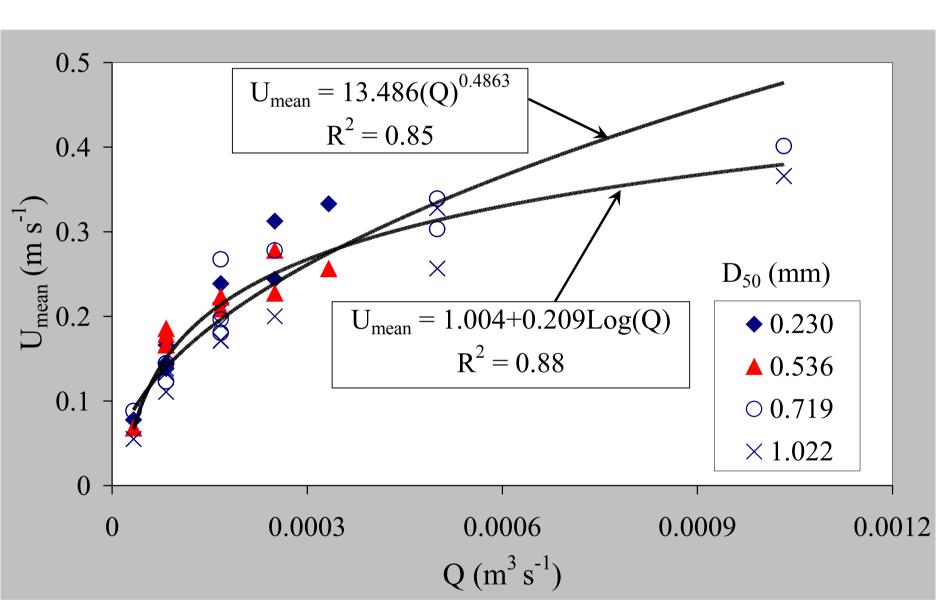
- L=3.0 m long and w=0.5 m wide rectangular (Fig. 1)
- Controlling variables:

 - 33 1033 * 10⁻⁶ m³ s⁻¹ Flow rates (Q): Substrate: Sand (D₅₀) 0.230, 0.536, 0.719 and 1.022 mm 3°, 5°, 7.5° and 10° Slopes (S):

Figure. 3a - 3d: Relationship between U_{depth} and U_{dye} for different substrates (Upper left to lower right).

Decreasing effect of grain size (Figure 3a – 3d, Table 1)

Table 1: Values for α depending on grain size.		
Grain size [mm]	α	
0.230	0.44	
0.536	0.77	
0.719	0.82	
1.022	0.82	



- Good fit of logaritmic function depending for relation of U_{mean} with Q
- Trends towards maxi-mum flow velocity U_{mean}

Measurements and Data Analysis

- Flow velocity:
 - Lycopene dye tracing over a test length of 1.24 m, 5 replicates
- Average water depth (h):
 - 2 point gauges, 0.1 mm resolution (Fig. 1, right)
- Mean flow velocity (U_{mean}) estimated according to:

 $U_{mean} \approx U_{depth} =$

- Bed form gathered by laser scanning (Fig. 1)
- Regression analysis with PASW between T_c and controlling variables or composite force predictors

Results and Discussion

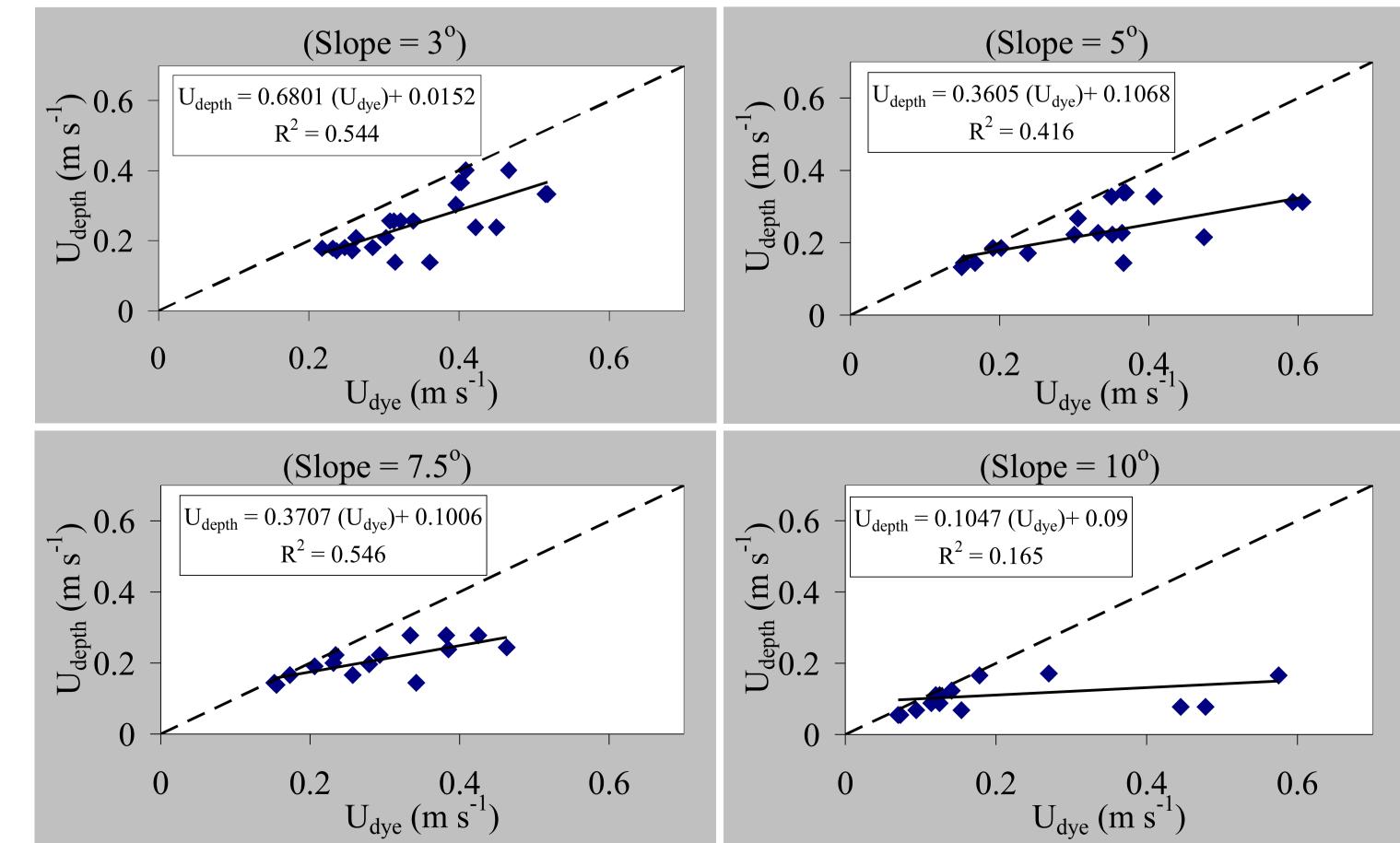
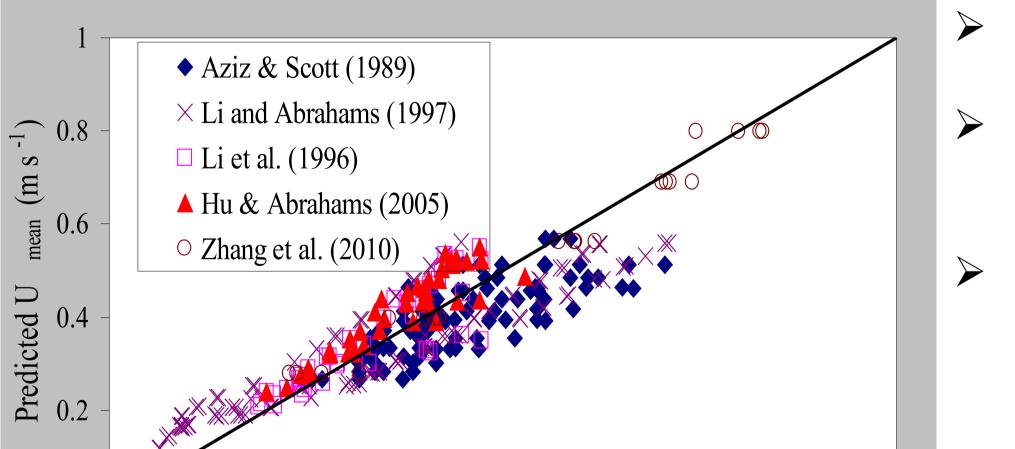


Figure 4: Relation between U_{mean} and Q.

Table 2: Results from stepwise regression analysis for U_{mean} depending on Q, D_{50} and S

Eq. No.	Equation	R ²
1	$U_{_{mean}}=13.486(Q)^{_{0.4863}}$	0.85
2	$U_{mean} = 1.004 + 0.209 Log(Q)$	0.88
3	$Log(U_{mean}) = 0.645 + 0.506 Log(Q) - 0.172 Log(D_{50})$	0.89
4	$U_{mean} = 1.072 + 0.217 Log(Q) - 57.69(D_{50})$	0.92
5	$Log(U_{mean}) = 0.646 + 0.502 Log(Q) - 0.17 Log(D_{50}) - 0.01 \log(S)$	0.89
6	$U_{mean} = 1.085 + 0.222 Log(Q) - 59.21(D_{50}) + 0.001(S)$	0.92



- functions Best Of fit including all variables
- Influence S of verv small and can be neglected
- Validation with different datasets from literature shows best fit with eq. 3

Figure 2a - 2d: Relationship between U_{depth} and U_{dye} for different slopes. (Upper left to lower right)

- Dye velocity (U_{dye}) always higher than estimated mean velocity (U_{mean})
- Effect highest for steep slopes and high velocities
- Effect of slope weak, no correlation with steep slopes

0.2 0.4 0.8 0.6 Measured U_{mean} (m s⁻¹)

Figure 5: Application of eq. 3 to different data sets of literature.

5 Conclusions

- No constant relationship between dye tracer velocity (U_{dye}) and mean velocity (U_{mean})
- Clear impact of flow rate (Q) and grain size (D_{50}) on flow velocity
- In mobile beds, slope (S) has only a weak impact on flow velocity, especially as U_{mean} tends to reach a maximum value with increasing Q
- Reasons have to be found in complex interaction of sediment uptake, changes in flow characteristics etc.

References

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